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## Method development for measuring <sup>17</sup>O-excess on the WAIS Divide Ice Core

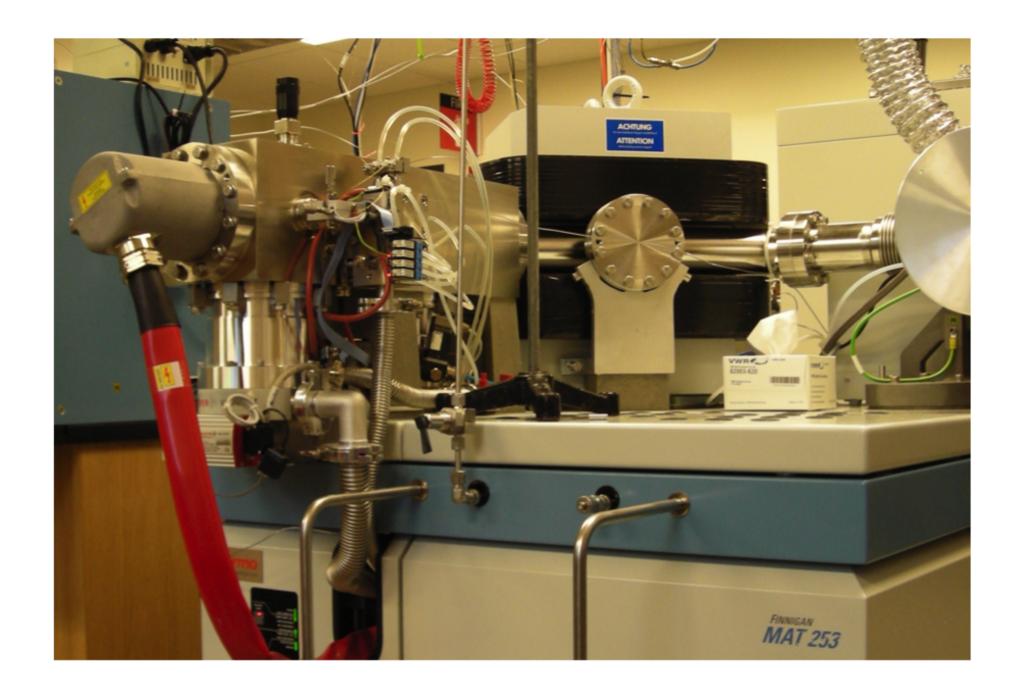


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#### Introduction

Stable isotope ratios of water in polar precipitation, as measured in ice cores, have been particularly useful in the interpretation of past climates. While  $\delta^{18}O$ alone is used as a proxy of past temperature, the recent development of techniques to measure the <sup>17</sup>O/<sup>16</sup>O ratio precisely (Barkan & Luz, 2005) has allowed <sup>17</sup>O-excess to be added to the ice-core isotope toolbox. The combination of  $\delta^{18}O$ ,  $\delta^{17}$ O and  $\delta D$  – giving the parameters <sup>17</sup>Oexcess and deuterium excess – could provide valuable additional information on the evaporative conditions of the oceanic moisture sources for Antarctic precipitation. Deuterium excess has been widely used, but with mixed results due to its sensitivity to conditions both during evaporation and during precipitation; this is theoretically less problematic for <sup>17</sup>O-excess.

The  $O_2$  sample is then run on a ThermoFinnigan MAT 253 isotope ratio mass spectrometer (IRMS). The O2 sample is analyzed for mass 32, 33, and 34 abundance measurements for  $\delta^{18}O$  and  $\delta^{17}O$ in dual-inlet mode by multicollector mass spectrometry.



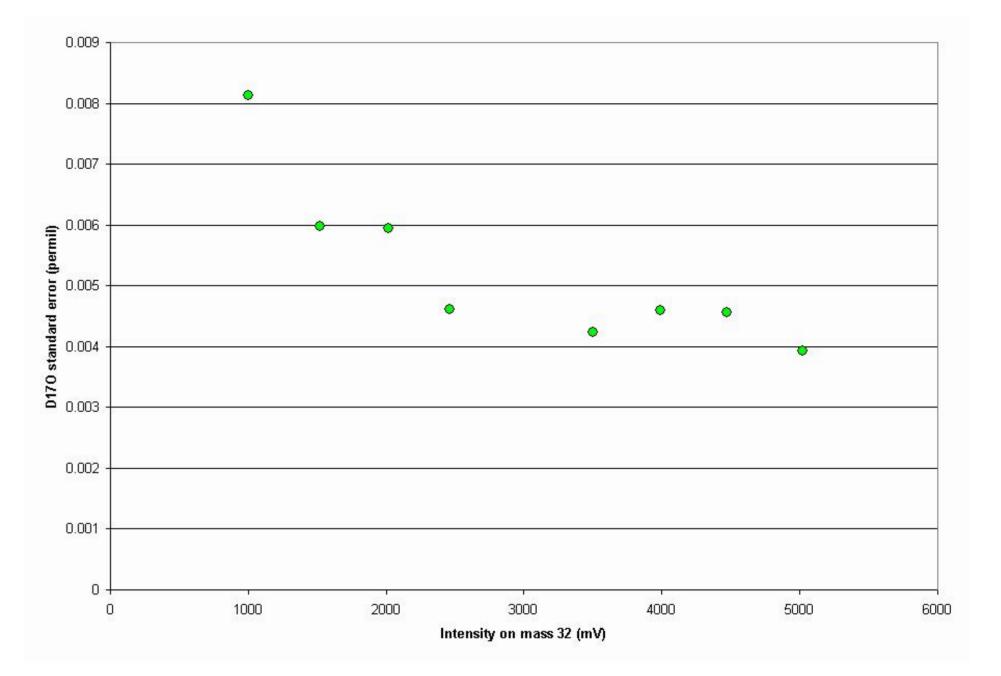


Figure 4. Standard error of  $\Delta^{17}$ O from high-purity  $O_2$  reference gas for a range of mass 32 intensities (mV) from a total of 80 acquisitions (run). The mass spectrometer error (standard error of the mean) for  $\Delta^{17}$ O has reached 0.004‰, (4 per meg) for higher intensities.

## Definition

<sup>17</sup>**O-excess** =  $\delta'^{17}$ O – 0.528 x  $\delta'^{18}$ O, where  $\delta'^{18}$ O = ln(R<sub>SAM</sub>/R<sub>SMOW</sub>) where R<sub>SAM</sub> =  $^{\times}$ O/ $^{16}$ O in sample, R<sub>SMOW</sub> =  $^{\times}$ O/ $^{16}$ O in Vienna Standard Mean Ocean Water. x = 18, 17. Note the different definition of 'delta' from the conventional one:  $\delta^{\times}$ O = R<sub>SAM</sub>/R<sub>SMOW</sub> -1. (Barkan & Luz, 2005) Figure 2. ThermoFinnigan MAT 253 Isotope Ratio Mass Spectrometer (IRMS) with dual inlet. O<sub>2</sub> sample is collected in stainless steel cold finger and then placed vertically in sample side of dual inlet.

Each mass spectrometric measurement is comprised of three separate acquisitions during which the ratio of sample to reference is determined 25 times. Each of theses cycles consists of 26 seconds of integration to maximize counting statistics and 30 seconds of idle time between each cycle to eliminate any residence time (memory) effects. The pressures of the sample and reference gas are balanced before each acquisition and the reference gas is refilled before each set of three acquisitions.

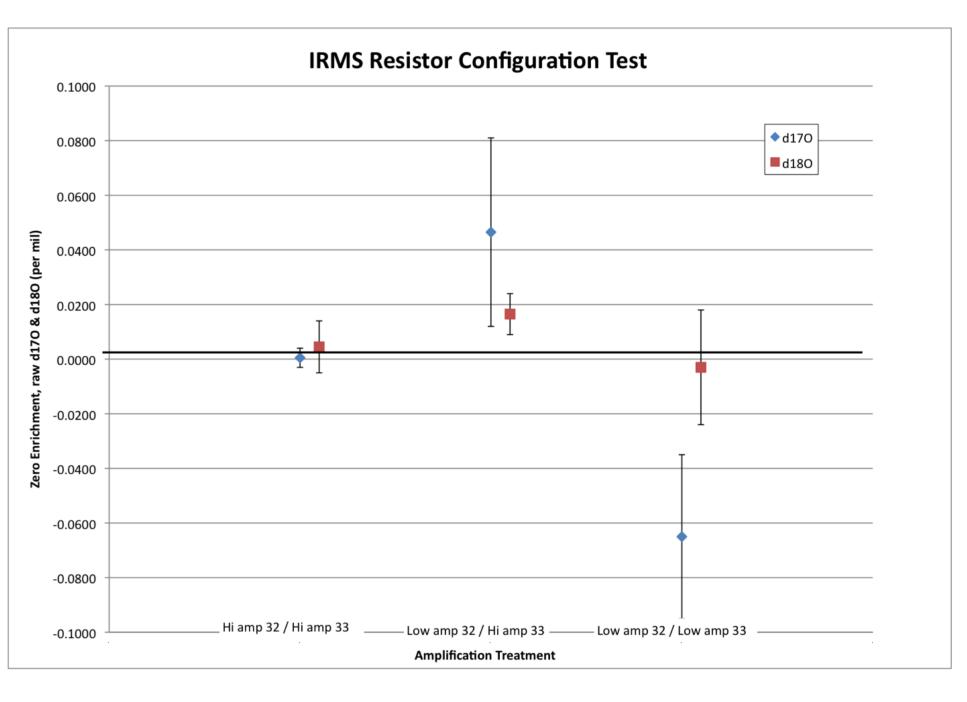


Figure 5. Test of resistor configuration. Note that high amplification treatment on both Mz 32 & 33 produces the highest precision.

#### Summary

## **Materials & Methods**

The analytical method for determination of the oxygen isotopic ratios is based on work by Barkan & Luz, 2005. To start, 5  $\mu$ L of water are injected via syringe through a high temp (400 °C) septum into a CoF<sub>3</sub> reaction tube, instantaneously converting H<sub>2</sub>O to O<sub>2</sub> and creating HF and Co as a byproduct. The O<sub>2</sub> sample is cyrofocused in a liquid nitrogen (LN<sub>2</sub>) trap, and then transferred to a cold finger in a liquid helium (LHe) dewar. O<sub>2</sub> samples in stainless steel cold fingers are warmed to room temperature and inserted vertically, one at a time, into the dual inlet port of the IRMS.

# Water Fluorination Line O2 Purification H2O Conversion to O2 Helium Purification

#### Results

Refinement of mass spectrometer integration period (seconds) and intensity range (mV) has improved precision of the  $\Delta^{17}$ O measurement to a standard error of 4 per meg. WAIS water standards put through the <sup>17</sup>O Water Fluorination line have been run on the IRMS, but with poor reproducibility thus far. Currently, a small (1-2 per mil) decline of isotopic values occurs during the entire length (25 cycles) of the run. The primary cause for this trend is most likely a decrease of the sample gas volume and pressure within the sample bellows. Measuring <sup>17</sup>O-excess requires exceptionally high precision (per meg) and sample reproducibility. We are continuing to pursue the shot noise limits of the mass spectrometer to achieve the best precision we can and to hone the <sup>17</sup>O Water Fluorination process for consistent reproducibility of WAIS water standards, before running actual ice-to-water samples

#### References

- Abe, O. (2008), Isotope fractionation of molecular oxygen during adsorption/desorption by molecular sieve zeolite, Rapid Communications in Mass Spectrometry, 22, 2510-2514. Baker, L., I. Franchi, J. Maynard, I. Wright, and C. Pillinger (2002), A technique for the determination of 180/160 and 170/160 isotopic ratios in water from small liquid and solid samples, Anal. Chem, 74, 1665-1673. Barkan, E., and B. Luz (2003), High-precision measurements of 170/160 and 180/160 of 02 and 02/Ar ratio in air, *Rapid Communications in Mass* Spectrometry, 17, 2809-2814. Barkan, E., and B. Luz (2007a), Diffusivity fractionations of H<sub>2</sub><sup>16</sup>O/H<sub>2</sub><sup>17</sup>O and  $H_2^{16}O/H_2^{18}O$  in air and their implications for isotope hydrology, *Rapid* Commun. Mass Spectrom, 21, 2999-3005. Landais, A., E. Barkan, and B. Luz (2008), Reply to comment by Martin F. Miller on "Record of d180 and 170-excess in ice from Vostok Antarctica during the last 150,000 years," Geophys. Res. Lett. Luz, B., E. Barkan, R. Yam, and A. Shemesh (2009), Fractionation of oxygen and hydrogen isotopes in evaporating water, *Geochimica Et Cosmochimica Acta*.
- Risi, C., A. Landais, S. Bony, J. Jouzel, and V. Masson- (2010), Understanding the 170 excess glacial-interglacial variations in Vostok precipitation, *Journal of*

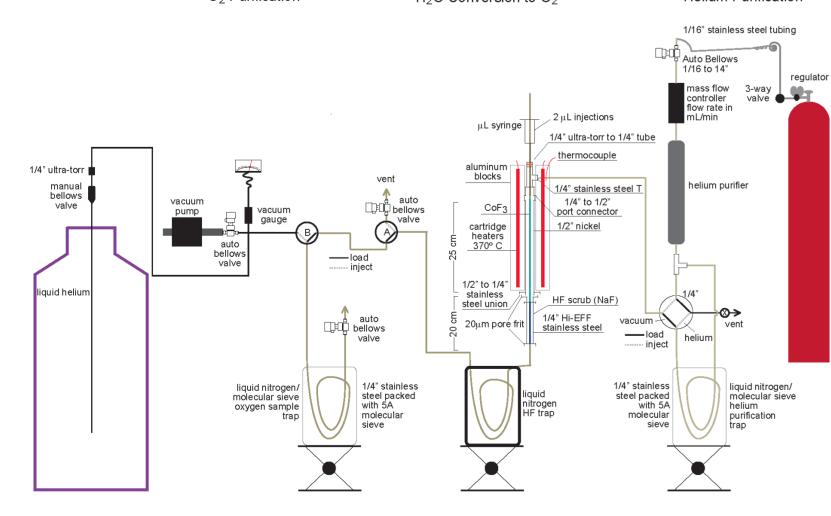


Figure 1. Plumbing schematic of <sup>17</sup>O Water Fluorination system showing He flow, injection port,  $CoF_3$  reaction heater,  $LN_2$  traps, and LHe dewar. Fluorination of water converts H<sub>2</sub>O to O<sub>2</sub>, which is purified by NaF scrub, and cyrofocused using a molecular sieve trap held at liquid nitrogen temperature. The O<sub>2</sub> sample is then transferred to a cold finger in a liquid helium dewar.

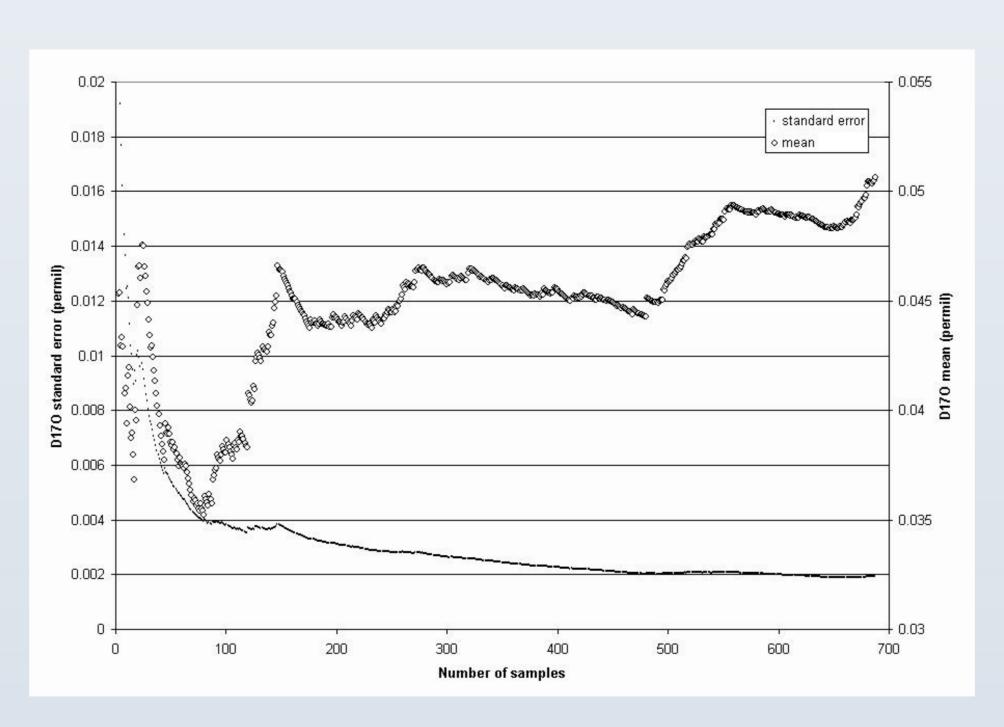
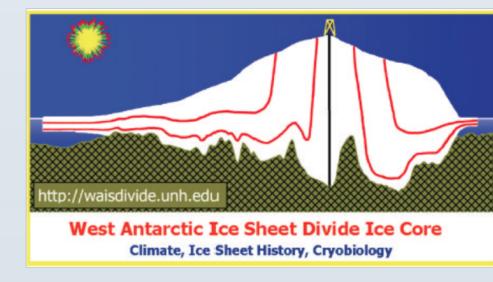


Figure 3. Standard error of  $\Delta^{17}$ O approaches (2 per meg) with increasing measurement number. Mean  $\Delta^{17}$ O shows possible drift after 500 samples. In this plot, each point represents one comparison of sample and standard; consisting of 26 seconds of integration and 30 seconds of idle time between each cycle. Geophysical.

Uemura, R., O. Abe, and H. Motoyama (2010), Determining the 170/160 ratio of water using a water-CO2 equilibration method: application to glacialinterglacial changes in 170-excess from the Dome Fuji ice core, *Geochimica Et Cosmochimica Acta*.

#### Acknowledgements

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**Further Information** 

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